

Discussion

Comment on “Characteristics and dynamic origin of the large-scale Jiaoluotage ductile compressional zone in the eastern Tianshan Mountains, China” by X.W. Xu, T.L. Ma, L.Q. Sun and X.P. Cai [Journal of Structural Geology 25 (2003) 1901–1916][☆]

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1. Introduction

In a recent paper, Xu et al. (2003) presented their study on the Jiaoluotage ductile compressional zone (JDCZ) in the eastern Tianshan Mountains, China. Xu et al. (2003) proposed a model of the geodynamic evolution of the JDCZ and stated that during the Carboniferous, the Tarim paleo-ocean plate was subducted beneath the Jungar plate, resulting in the separation of the middle Tianshan terrane from the Jungar plate and the formation of the middle Tianshan arc and Jiaoluotage back-arc basin. Although we do not dispute their main conclusions related to the ductile deformation, we disagree with Xu et al.'s model of the geodynamic evolution of the JDCZ. In this paper, we will discuss the following issues:

- (1) Was the central Tianshan (named the middle Tianshan by Xu et al. (2003)) separated from the Jungar plate?
- (2) When and how did the central Tianshan arc form?
- (3) Was Jiaoluotage a back-arc basin during the Carboniferous?
- (4) The model of the geodynamic evolution of the JDCZ.

2. Was the central Tianshan separated from the Jungar plate?

Up to now, no Precambrian rocks have been observed in

or around the Jungar basin (Allen et al., 1993b; Charvet et al., 2002; Li et al., 2000). Only the Paleozoic volcanics, volcanoclastics and granitoids crop out in the northern margin of the Tianshan and the eastern Jungar (Feng et al., 1989; Carroll et al., 1990; Laurent et al., 2003). In contrast, the Precambrian basement rocks and Sinian tillites occur widely in the central Tianshan and the northern margins of the Tarim. In the eastern segment of the central Tianshan, Precambrian orthogneisses, migmatites, schists, amphibolites, schists and marbles are exposed in northern Kumux–Maanqiao, Tucileike, Weiya and Xingxingxia areas (BGMRX, 1993; Fig. 1). In the western segment, Precambrian amphibolite facies metamorphic rocks, carbonates, clastic rocks and minor tillites crop out predominantly in Wenquan, Nalati, Baluntai, Borohoroshan and Haxiliagen-daban–Bingdaban regions (Gao et al., 1998). On the northern margins of Tarim, the Sinian (Vendian or latest Proterozoic) neritic–littoral clastic rocks, carbonates and tillites were also observed (Ma et al., 1993), which lie disconformably on the Precambrian basement (called the Aksu Group) in the Kelpin region (Nakajima et al., 1990). From the Cambrian to Early Ordovician, passive continental margin sediments were deposited along both the northern margin of the Tarim block and the central Tianshan belt, which were characterized by carbonate rocks, cherts and terrigenous coarse clastic rocks. There are no Carboniferous volcanic rocks in the eastern segment of the central Tianshan. These facts demonstrate that the Tarim block and the central Tianshan might have been consolidated before the Ordovician. So far, there is no evidence to support the conclusions by Xu et al. (2003). We suggest that a southward-dipping oceanic subduction took place from

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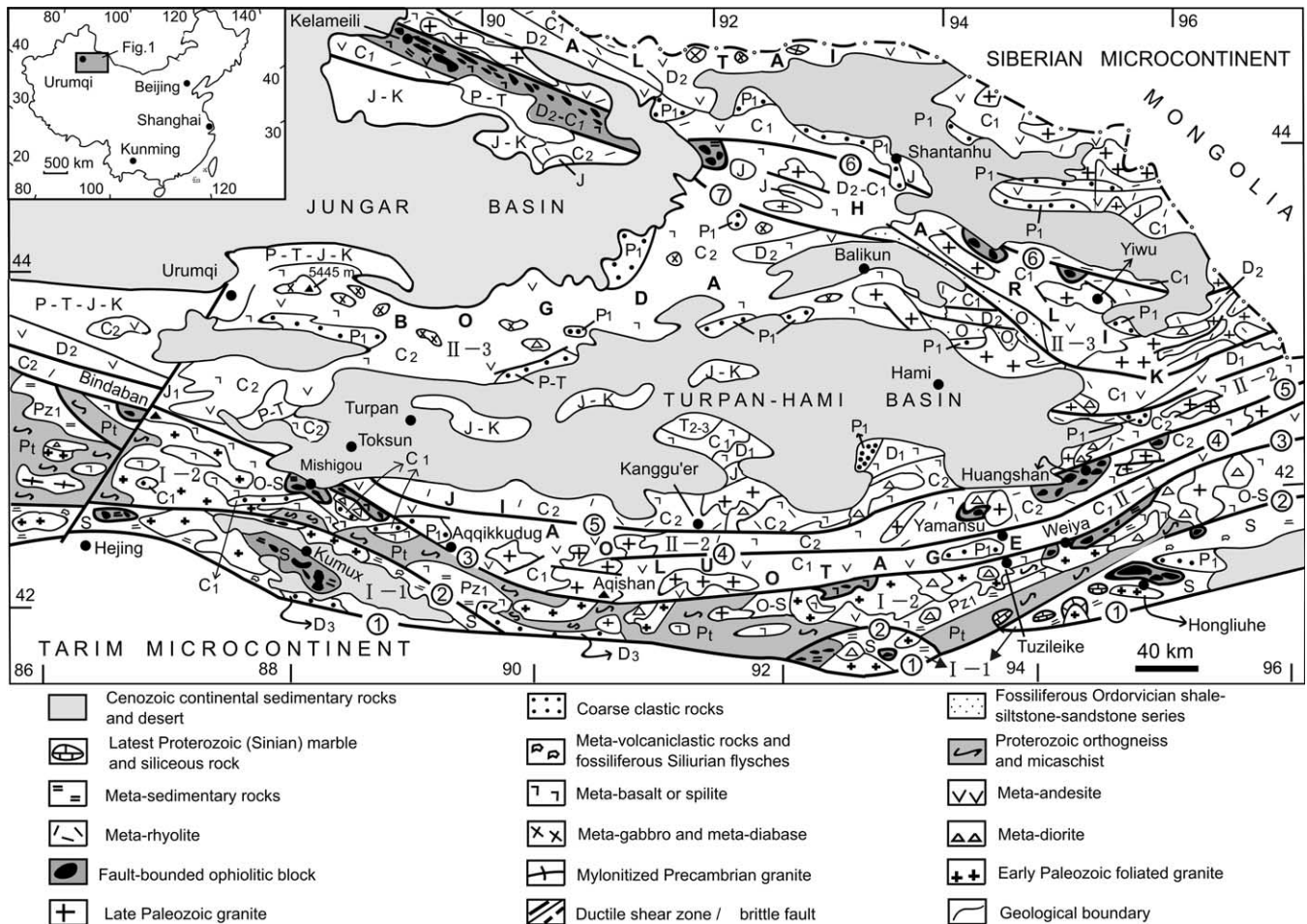


Fig. 1. Simplified geological map of the east Tianshan range (modified from Shu et al. (2002)). ① northern marginal fault of Tarim; ② southern marginal fault in central Tianshan; ③ Aqqikkudug–Weiya suture zone; ④ Aqishan–Yamansu fault; ⑤ Kangguer–Huangshan fault; ⑥ Karamaili–Yiwu suture zone; I–1 Kumux–Hongliuhe suture zone; I–2 central Tianshan Early Paleozoic volcanic arc with the Precambrian basement; II–1 Aqishan–Yamansu Devonian–Carboniferous volcanic arc; II–2 Kangguer–Huanshan Carboniferous inter-arc basin; II–3 Bogda–Harlik Devonian–Carboniferous volcanic arc.

the Middle Ordovician to the Middle Silurian along the Aqqikkudug–Weiya zone (named North Tianshan Suture Zone by Laurent et al. (2003)), forming the central Tianshan arc and the southern Tianshan back-arc basin that led to the separation of the central Tianshan from the Tarim block (Ma et al., 1993; Zhu et al., 1993; Gao et al., 1998; Charvet et al., 2002).

3. When and how did the central Tianshan arc form?

The Middle Ordovician basalt, andesite, dacite, rhyolite, greywacke and Silurian flysch crop out in the central Tianshan. The Ordovician–Silurian meta-volcanic sedimentary rocks are in contact with the Precambrian basement by ductile faults (Shu et al., 2002). The diorite in Tucileike (Fig. 1) gives a Rb–Sr whole-rock age of 460 ± 11 Ma (Yang, 1988) and the andesite in southern Aqishan yields a Rb–Sr isochron age of 477 Ma (Lu, unpublished data, in: Ma et al., 1993). Early Carboniferous red molasses, sandstone, mudstone, marlite and limestone are

unconformably overlaid on the Ordovician meta-volcanic rocks or Silurian flysch layer (Charvet et al., 2001; Shu et al., 2002). No Carboniferous volcanic rocks occur in the eastern segment of the central Tianshan, unlike Xu et al.'s depiction (Xu et al. (2003) and their Fig. 18). Xu et al. (2003) proposed that a northward-dipping oceanic subduction along the southern margin of the central Tianshan created the late Paleozoic central Tianshan arc during the Carboniferous. Kinematic and chronological research, however, suggests that a southward-dipping oceanic subduction took place along the Aqqikkudug–Weiya zone during the Ordovician and produced the early Paleozoic central Tianshan arc (Ma et al., 1993, Shu et al., 2000).

4. Was Jiaoluotage a back-arc basin during the Carboniferous?

We do not agree with Xu et al.'s (2003) proposal that the Jiaoluotage is a back-arc basin resulting from a northward-dipping subduction of the Tarim paleo-ocean plate during

the Carboniferous. Xu et al. (2003) did not present the Paleozoic rock assemblages of Jiaoluotage. In fact, two Paleozoic tectonic units were recognized in Jiaoluotage. (1) Aqishan–Yamansu unit: the Devonian–Late Carboniferous island-arc type volcanic and volcanoclastic rocks are predominantly exposed in the Aqishan–Yamansu and Dananhu areas. They are mainly andesite, dacite, dacitic rhyolite, rhyolite, rhyodacite, andesitic breccia and andestic tuff, and have been dated by whole-rock Rb–Sr isochron at 290 ± 5 Ma for andesite and 300 ± 13 Ma for rhyolite (Li et al., 1998) and by whole-rock Sm–Nd isochron at 416 ± 120 Ma for basalt (Rui et al., 2002). The chemical composition and the REE distribution patterns of the volcanic rocks show that they are calc–alkaline (Ma et al., 1993). (2) Kangguer–Huangshan unit: this unit lies between

the Aqishan–Yamansu and the Bogda–Harlik Devonian–Carboniferous volcanic arcs. This Middle Carboniferous inter-arc basin is evidenced by the following geological facts: (a) the Early Carboniferous sedimentary facies in the Kangguer–Huangshan unit exhibit an environmental evolution from shallow water (fluvial) to deep water (flysch) containing radiolarian cherts, which reveals the spreading and extension of the basin (Guo and Ma, 1990); (b) the Middle Carboniferous bimodal volcanic rocks exposed in the Kangguer unit indicate a rifting environment (Zhu et al., 1996); (c) a large Alaskan-type mafic–ultramafic complex was emplaced along the Kangguer fault (Ma et al., 1993; Zhu et al., 1996; Charvet et al., 2002), and was dated by Sm–Nd isochron at 308.9 ± 10.7 Ma (Li et al., 1998). These geological data indicate that Kangguer–Huangshan unit

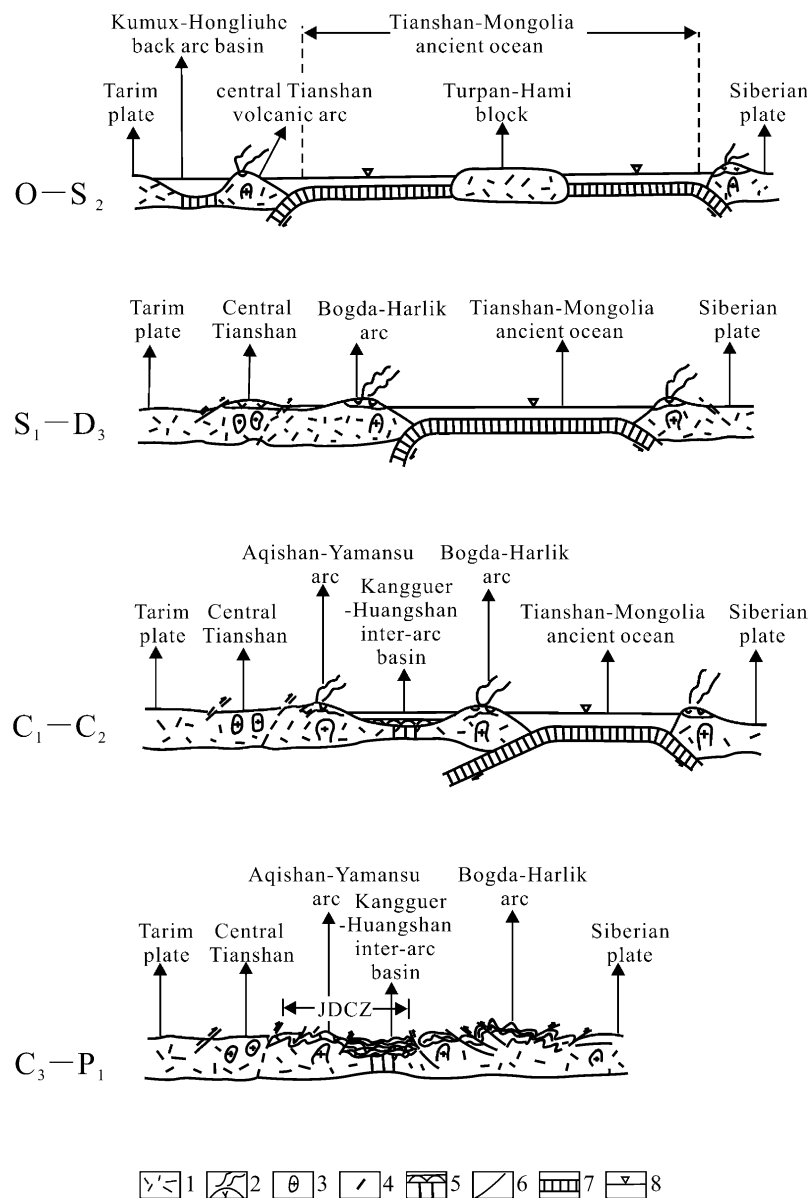


Fig. 2. Model of Paleozoic tectonic evolution of east Tianshan: (1) continental crust; (2) island arc; (3) granites; (4) ophiolite; (5) mafic–ultramafic rocks; (6) fault; (7) oceanic crust; (8) sea surface.

represents an inter-arc basin between Aqishan–Yamansu and Bodga–Harlik island-arcs.

5. The model of the geodynamic evolution of the JDCZ

In the Precambrian, a huge continent, which included Tarim, Yili and Turpan–Hami blocks, already existed in central Asia (Carroll et al., 1990; Ma et al., 1993; Charvet et al., 2001, 2002; Shu et al., 2002, 2003). During the early Cambrian, the E–W-trending Tianshan–Mongolia ancient ocean formed as a result of the rifting between the Tarim and the Yili–Turpan–Hami microcontinents (Shu et al., 2002). Middle Ordovician, southward-dipping oceanic subduction along the Aqikkudug–Weiya zone (Fig. 2) formed the Early Paleozoic island-arc-type volcanic rocks and granitoids in the central Tianshan (Charvet et al., 2001; Guo et al., 2002; Shu et al., 2002, 2003). As the Tianshan–Mongolia ancient ocean was subducted, the continental crust on the south of the central Tianshan began to extend, creating the Kumux–Hongliuhe marginal sea that separated the central Tianshan from the Tarim (Zhu et al., 1993; Gao et al., 1998; Charvet et al., 2001; Shu et al., 2002). Throughout the Middle Devonian–Middle Carboniferous, the new subduction zone occurred on the north side of the Turpan–Hami block. The Tianshan–Mongolia ocean subducted towards the south along the Karamaili–Yiwu zone, forming the Aqishan–Yamansu, Dananhu and Bogda–Harlik Devonian–Carboniferous volcanic arcs, and then the Kangguer–Huangshan Middle Carboniferous inter-arc basin in the Jiaoluotage region (Figs. 1 and 2). These three tectonic units were named the North Tianshan arc by Windley et al. (1990) and Allen et al. (1993a). During the late Late Paleozoic, collision took place between the northern Tianshan and the southern Siberian active continental margins, resulting in the accretion of the North Tianshan arc towards the central Tianshan and the closure of the Kangguer–Huangshan inter-arc basin, which formed the JDCZ (Fig. 2).

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